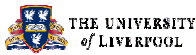


Propagation of wave signals along the western boundary and their link to ocean overturning in the North Atlantic



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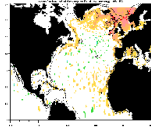


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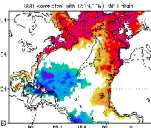
1. How changes in the high latitude forcing communicate over the ocean?

Downstream response at lower latitudes along the western boundary involving:

- rapid response: wave propagation against the sidewalls and along the equator – Kelvin waves, coastal trapped waves (Huthnance, 1978, JPO; Kawase, 1987; Johnson & Marshall, 2002, JPO)
- an intermediate response involving changes in local circulation
- slower response: advection along western boundary and evolution of dense water masses



Correlation of altimetry (Hughes & Meredith)



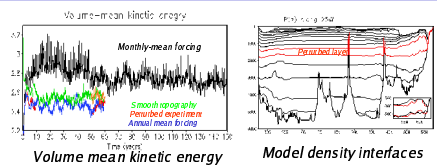
Model SSH correlation

2. Model Studies

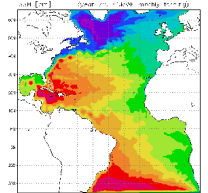
Model simulations carried out using a 16 layer isopycnic model (MICOM).

Resolution is 0.23° (26 km at the equator & 13km at 60°N). Model initialised from Levitus and run for 150 years, forced by ECMWF monthly-mean winds and surface fluxes. Parallel run is performed, forced by annual mean forcing, and a twin experiment involving extra thermohaline forcing is run from year 50. The twin perturbation experiments is run for 10 years, with deep interface raised 50m per 5 days over the northern relaxation zone.

Advection is monitored by a transient model tracer, released in the Labrador Sea, when the perturbed buoyancy forcing starts.



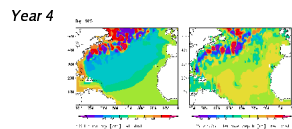
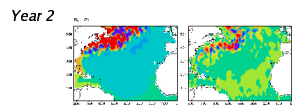
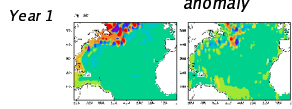
Model density interfaces along 25°W .



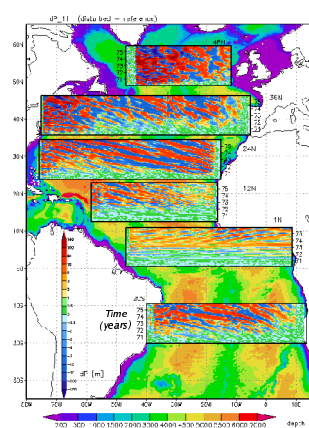
Model Sea-surface height

3. Propagation of wave signals anomalies in the model

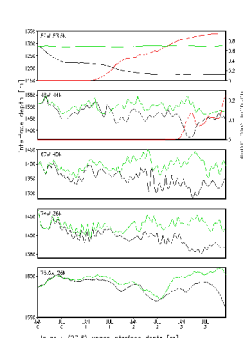
Model snapshots of: SSH Interface depth anomaly



Hovmoeller diagrams of Interface depth anomaly



Propagation of signals along the western boundary



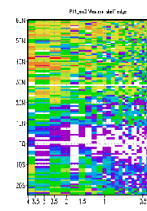
--- default; --- perturbed run; --- tracer signal.

- there is a rapid spreading along the western boundary
- a slower spreading along the eastern boundary, generating Rossby waves
- the SSH and isopycnal height anomalies appear related, but with SSH signals on a slightly broader scale.

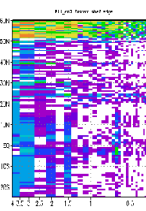
4. Spectral changes in boundary waves

Pressure time series spectral analysis

Western Atlantic



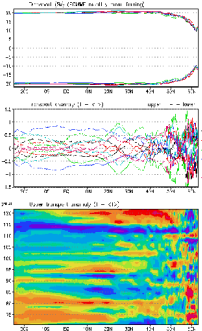
Eastern Atlantic



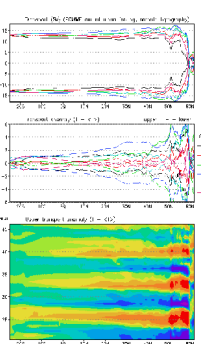
- higher frequency waves possibly filtered by the boundary current
- waves modified by the topography and stratification
- interaction with incoming Rossby waves
- lower frequencies and amplitudes along eastern boundary and low latitudes

5. How do the waves affect the overturning variability

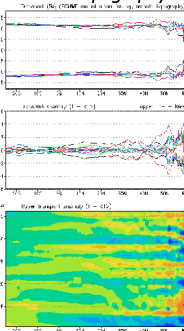
Realistic monthly-mean perturbations



Annual mean forcing



Annual mean forcing, smooth topography



Upper and deep transports

Transport anomalies

- transport variations of 1 – 2 Sv in mid-latitudes and 2 – 4 Sv at high latitudes
- fast changes in the overturning (few years after the high latitude perturbation)
- longer term transport oscillations

6. Conclusions

- Boundary wave propagation on time scale of months to years
- Connected to interior via equatorial Kelvin waves and basin-wide Rossby waves
- High frequency boundary waves are possibly filtered by the western boundary current
- Waves modified by the topography and stratification
- Wave signals associated with changes in basin-wide overturning of typically 1 – 3 Sv, occurring prior any deep advective signal

Poster available on-line: www.liv.ac.uk/~vassil/rapid